Monitoring Global Land Surface Drought Based on a Hybrid Evapotranspiration Model

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Objectives

- I. Introduce a linear combination of the most important parameters (air temperature, net radiation, etc.) controlling the actual ET to simplify the evaporative fraction (EF), and to then develop a new hybrid ET model. Also, to validate the hybrid ET model using Atmosphere Radiation Measurement (ARM), FLUXNET and data from Chinese experiments
- II. Use the evaporative drought index (EDI) to infer global land surface dryness conditions from 1984 to 2002 using the hybrid ET model with GEWEX, AVHRR-GIMMS-NDVI, and NCEP-2 products. The EDI patterns are compared with Palmer Drought Severity Index (PDSI) products to determine the validity of the method's drought assessment on a global scale.

Data

- Global monthly surface SW ↑↓ and LW ↑↓, derived from the Global Energy and Water
 Cycle Experiment (GEWEX), 1 deg resolution
- Monthly NDVI obtained from NOAA/AVHRR observations
- Monthly mean air temperature and T_{max} and T_{min} obtained from NCEP-2 reanalysis data

Data - Validation

- NDVI (modis) 16 day composites interpolated to daily values
- Flux tower data 12 ARM towers, 7 Fluxnet towers, 3 towers in China
- Palmer's drought index (1984 2000), 2.5 deg resolution

- Assume landscape is a mixture of bare soil and vegetation elements
- Proportion of vegetation obtained from vegetation fraction
- Total ET in a pixel = combination of ET from soil and ET from vegetation

$$ET = (1 - f_{veg})ET_{soil} + f_{veg}ET_{veg}$$
$$f_{veg} = \frac{NDVI - NDVI_{min}}{NDVI_{max} - NDVI_{min}},$$

NDVImin = 0.05 NDVImax = 0.95

• Introduce evaporation fraction:

$$EF = \frac{ET}{Q} = \frac{ET}{R_n - G} = \frac{ET}{(1 - G_0)R_n}$$

• Can use EF to describe both ETsoil and ETveg

$$ET_{soil} = Q_{soil}EF_{soil} = a_0Q * EF_{soil} = a_0(1 - G_0)R_nEF_{soil}$$
$$ET_{veg} = Q_{veg}EF_{veg} = a_1Q * EF_{veg} = a_1(1 - G_0)R_nEF_{veg}$$

• Evaporation fraction for soil, by Nishida et al., 2003:

$$EF_{soil} = \frac{Q_{soil0}}{Q_{soil}} \frac{T_{soil,\max} - T_{soil}}{T_{soil,\max} - T_a} = \frac{Q_{soil0}}{Q_{soil}} \left(1 + \frac{T_a - T_{soil}}{T_{soil,\max} - T_a}\right)$$

Assume that Q_{soil0}/Q_{soil} and (T_a –T_{soil})are all invariant constants. To calculate EF_{soil} approximately using air temperature, can replace (T_{soil,max} –T_a)with the diurnal air temperature range (T_{max} –T_{min}) to simplify the EF_{soil} by adding the empirical coefficients.

$$EF_{soil} = b_0 + \frac{b_1}{T_{\max} - T_{\min}}$$

Tmax, Tmin = daily maximum and minimum air temperatures b0, b1 = empirical coefficients

Evaporation fraction for vegetation by Nishida et al.,
2003:

$$EF_{veg} = \frac{\alpha\Delta}{\Delta + \gamma(1 + r_c/2r_a)}$$
$$\frac{1}{r_c} = \frac{f_1(T_a)f_2(VPD)f_3(PAR)f_4(\Psi)f_5(CO_2)}{r_{c_{-}\min}} + \frac{1}{r_{cuticle}}$$

 α = 1.26 Δ = desat/dT VPD = vappr pressure deficit Phi = soil water potential ra = aerodynamic resistance rc = canopy resistance

- <u>Simplifications</u>: develop a linear combination of WS, PAR, Tair, VPD, atmospheric CO₂ concentration and root soil moisture to approximately simulate (*).
 - WS, VPD and CO₂ not available from satellite omitted
 - Rn replaces PAR due to their approximately linear relationship
 - 1/(Tmax Tmin) selected to denote soil moisture near the soil

*

•
$$EF_{veg} = b_2 + b_3 + b_4 R_n + b_5 / (T_{max} - T_{min})$$

• Combination of all previous equations yields:

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$$ET = \left(1 - \frac{NDVI - 0.05}{0.95 - 0.05}\right) a_0(1 - G_0)R_n \left(b_0 + \frac{b_1}{T_{\text{max}} - T_{\text{min}}}\right) \\ + \frac{NDVI - 0.05}{0.95 - 0.05} a_1(1 - G_0)R_n \left(b_2 + b_3T_a + b_4R_n + \frac{b_5}{T_{\text{max}} - T_{\text{min}}}\right)$$

 Further simplification by integrating a series of empirical coefficients in order to propose a hybrid regression equation based on (**):

$$ET = R_n^2(c_1 NDVI - c_2) + R_n \left(c_3 + c_4 T_a + \frac{c_5}{T_{\text{max}} - T_{\text{min}}} \right) + R_n \left(c_6 + c_7 T_a + \frac{c_8}{T_{\text{max}} - T_{\text{min}}} \right) NDVI$$

Estimation of potential evapotranspiration (PET)

- PET represents the ideal evaporation rate for capturing the response to forcing variables if soil moisture is unlimited.
- Adopted the Hargreaves method to estimate PET this method is simple for practical use because it requires only two easily accessible parameters: temperature and solar energy

 $PET = 0.0023R_a(T_{mean} + 17.8)\sqrt{T_{max} - T_{min}}$

Model Validation

 Randomly divided the 22 flux towers into 2 groups – used group 1 to estimate the coefficients, used group 2 to validate the results

Results



Scatter plots of the predicted monthly ET and ground-measured ET for the first group of flux towers using the second group data to calibrate the eight parameters

Results



Scatter plots of the predicted monthly ET and ground-measured ET for the second group of flux towers using the first group data to calibrate the eight parameters in Eq

Global Implementation

$$ET = R_n^2 (0.00084NDVI - 0.000978) + R_n \left(0.3044 + 0.0029T_a + \frac{0.284}{T_{\text{max}} - T_{\text{min}}} \right) + R_n \left(0.1273 + 0.01T_a + \frac{0.065}{T_{\text{max}} - T_{\text{min}}} \right) NDVI$$

"Sufficiently representative" except for desert and glacier regions.

Global implementation - 1986



An example of the comparison of monthly ET using GEWEX, AVHRR-GIMMS-NDVI and NCEP-2 datasets of October 1986, and the corresponding latent heat flux from the GSWP-2 datasets. (a) The spatial distribution of the estimated monthly ET obtained using Eq. (21) for October 1986; (b) spatial distribution of the corresponding GSWP-2 ET; and (c) scatter plots of the estimated monthly ET for October 1986 and the corresponding GSWP-2 ET.



Fig. 6. Estimated global monthly ET for April–September of 2001–2002.

EDI vs. PDSI



Fig. 7. Temporal variation of monthly normalized EDI and PDSI anomalies from January 1984 to December 2002.